



FIGURE 2. Map of northeastern Ohio showing surface distribution of Illinoian drift and Wisconsin till units (Modified from White, 1967.)

At the pit of the Ravenna Sand and Gravel Company, about 1/4 mile south of State Route 5, on New Milford Road, Ravenna Township (Fig. 9), in Portage County, as much as 50 feet of Illinoian drift is exposed (Winslow and White, 1966). The lowest unit of the drift is a brown, calcareous, sandy clay till that is exposed in the floor of the pit, near its east end. The till contains an abundance of boulders, some of which are over one foot in diameter. Many of the boulders are striated and faceted as the result of abrasion while embedded in the base of the ice.

Overlying the till is 25-30 feet of gray gravel and brown sand, conspicuously cross-bedded in cut and fill structures. The high percentage of carbonate pebbles (32% limestone and dolostone) and crystalline rock fragments (23%), compare favorably with similar percentages found in Illinoian gravels at localities in Cuyahoga and Geauga counties (Winslow and White, 1966). A brown colored, thin, silty to sandy textured, oxidized till, as much as three feet thick, was formerly exposed in the pit, above the gravel and below a gray, coarse textured, calcareous, Wisconsin (Mogadore) till. The stratified sand and gravel between the two Illinoian till sheets may indicate an episode of ice retreat within the Illinoian stage of glaciation (Winslow and White, 1966).

SANGAMON

Physical evidence for the Sangamon interglacial stage is poorly preserved in the guidebook area. Erosion during the subsequent Wisconsin glacial stage probably removed much of the record. The weathered zones in Illinoian drift at the Garfield Heights and Ravenna Sand and Gravel Company Pits were developed during this interval. Fossil gastropods reported from the Sangamon soil at the Garfield Heights pit (Leonard, 1953) consist of several species which now reach the southern limits of their range far to the north in Canada. The fact that they were able to live in Ohio during the Sangamon suggests that summers were cooler than those now occurring in this area.

WISCONSIN

The Wisconsin ice advancing into northeastern Ohio was split into two lobes as a result of topography. The western segment, the Killbuck Lobe, moved southeast down the low area between the high land in Richland County and the Summit-Geauga County upland. To the east of the Summit-Geauga high land, the Grand River lobe advanced southward through the Grand River lowland (White, 1967). The location of these lobes is shown in Fig. 2.

The Wisconsin ice underwent a number of fluctuations in which the ice terminus retreated varying distances to the north. At least five re-advances of the ice are recorded in northeastern Ohio by till sheets of distinctive lithology (Fig. 2; Table 1).

In general, the age of the surface drift exposed at a given location in the guidebook area can be determined by comparison with Fig. 2 and Table 1. Exposures in some deeper gravel pits, strip mines, and stream valleys, however, may contain several superimposed drift sheets. The proper identification of these multiple till sheets usually requires mechanical analysis of the sand-silt-clay ratios and careful measurements of the depth of oxidation and leaching.

TABLE 1 — Properties of Wisconsin age glacial deposits in northeastern Ohio. (Based on data published in White, 1960, 1961, 1967; DeLong and White, 1963; Winslow and White, 1966; Goldthwait, et al., 1961.)

LOBE	UNIT	LITHOLOGIC TYPE	COMPOSITION, %			THICKNESS	OXIDIZED COLOR	DEPTH OF WEATHERING	
			SAND	SILT	CLAY			OXIDIZED and LEACHED	OXIDIZED
GRAND RIVER LOBE	ASHTABULA TILL	Silty clay till	23	48	29	0-50'+	Brown	3' 11"	5' 9"
	HIRAM TILL	Clay till	12	41	47	0-30'+	Dark brown	3'	12'
	LAVERY TILL	Silty clay till	24	45	31	0-15'	Dark brown	4' 7"	12' 7"
	KENT TILL	Silty sandy till	31	46	23	0-100'+	Yellow brown	5' 8"	10' 1"
	MOGADORE TILL	Sandy till	46	43	11	0-90'	Yellow- to olive-brown	7' 4"	12' 2"
KILLBUCK LOBE	ASHTABULA TILL	Silty clay till	?	?	?	0-20'	Brown	1' 5"	11"
	HIRAM TILL	Clay till	26	46	28	0-12'	Dark brown	2' 9"	4' 6"+
	HAYESVILLE TILL	Silty clay till	26	46	28	0-12'	Dark brown	3' 11"	2' 10"
	NAVARRE TILL	Silty sandy till	47	37	16	0-15'+	Yellow brown	5' 1"	3' 8"
	MILLBROOK TILL (ILLINOIAN?)	Sandy till	43	42	15	0-21'	Yellow- to olive-brown	9' 3"	13' 9"

Short intermittent retreats of the ice to the north during the Wisconsin glacial stage permitted re-establishment of normal surface drainage in the area. Some of the northward flowing streams were dammed by the ice terminus and/or moraines left in the valleys by the retreating ice. Water was ponded behind these dams, forming a number of different lakes that waxed and waned with the fluctuations of the ice terminus. Lake sediments deposited in these lakes can be seen along the valleys of many of the streams that drain north into Lake Erie. Although the presence of these deposits has been known for 80 years (Claypole, 1887), they have not been carefully studied. They are all probably Wisconsin in age, but their exact placement in the sequence of Wisconsin Age events is not known.

Fluctuations of the ice margin are also recorded in the guidebook area by deposits of wind-blown silt, or loess. Withdrawal of the ice left large areas of drift exposed at the surface without a cover of vegetation. The wind picked up the finer materials from the drift and re-deposited it as loess. Loess has been reported from below Wisconsin tills of different ages in northeastern Ohio (White and Totten, 1967). At Garfield Heights, in Cuyahoga County (Fig. 8), as much as 9 feet of silt (loess?) rests on a buried Sangamon soil and in turn is covered by lake sediments and the Hiram Till. Twenty species of snails have been reported from this deposit by Leonard (1953). The upper portion of this loess sequence has been dated by radiocarbon at $28,195 \pm 535$ years (White, 1965). The lower portion of the immediately overlying laminated silts (varves?) contains degraded plant fragments, insects (Coope, 1968; White, 1968) and mollusks. Wood from this unit has been dated at $24,600 \pm 800$ and $23,313 \pm 391$ years before the present (White, 1968).

ANCESTRAL GREAT LAKES

The last Wisconsin ice, represented by the Ashtabula Till (Fig. 2), withdrew from northeastern Ohio between 14,000 and 15,000 years ago. The present Great Lakes started their rather complex history about this period of time. Fluctuations in the water levels in these predecessors of the modern Great Lakes were controlled by oscillations in the margin of the late Wisconsin ice lobes; downcutting of some lake outlets by stream erosion; and regional crustal uplift due to the removal of the thick load of glacial ice. Most of the former lake levels in the Erie Basin are recorded in a series of beach ridges which occur above the present level of Lake Erie (Table 2; Fig. 1). Frank Carney mapped these features during the early part of the 20th century. Although these maps were never published, copies can still be purchased from the Ohio Geological Survey.

TABLE 2

Sequence of lake stages in the Erie Basin (modified from White, 1965)

Years before present	Lake stage	Elev. of shore
0	Erie	572'
6,000	low water	570' (?)
10,000	Lundy	620'
	Elkton	640'
	Warren III	675'
	Wayne	660'
12,000	Warren II	680'
	Warren I	690'
13,000	Whittlesey	735'
	Arkona	695', 700', 710'
14,000	Maumee III	785'
	Maumee II	760'
	Maumee I	790'

TABLE 1 - LOW-FLOW FREQUENCY INDICES

Station No. USGS	Station Name	Location	Drainage area in sq mi	Period of record water years	Mean Annual Discharge 1931-60 cfs	Discharge Exceeded 90% of Time cfs	Ratio of to mean annual discharge	2-Year Frequency 7-Day Duration	
								Average Low-Flow Discharge cfs	Ratio of to mean annual discharge
(Continued)									
38.8	4-1942	Tousaint Creek near Limestone	74.7	1939-63	648.4	419	0.95	0.196	0.0502
39	4-1946	Portage River near Pemberville	334	1931-35	426	487	0.30	0.0979	0.0093
40	4-1950	North Branch Portage River near Bowling Green	64.0	1928-32	436.9	466	0.67	0.172	0.105
41	4-1955	Portage River at Woodville	433	1929-31; 1940-63	4307	468	0.75	0.166	0.00418
42	4-1960	Sandusky River near Buyrus	89.8	1928-36; 1939-61	677.4	0.587	0.10	0.0871	0.0233
43	4-1965	Sandusky River near Upper Sandusky	299	1928-35; 1939-63	4236	510	0.40	0.0871	0.0178
46	4-1970	Sandusky River near Waterloo	776	1924-35; 1939-63	4587	464	0.80	0.066	0.012
47	4-1980	Sandusky River near Fremont	1,246	1924-35; 1939-63	4916	474	0.10	0.036	0.014
47.3	4-1980.1	Green Creek near Fremont	61.6	1929-36; 1961-63	653.6	504	0.15	0.28	0.112
47.8	4-1980.2	West Branch Huron River near Monroeville	286	1960-63	4188	538	0.20	0.049	0.036
50	4-1990	Huron River at Milan	343	1931-63	4297	529	0.10	0.0471	0.0343
51	4-1995	Vermilion River near Vermilion	260	1931-63	4258	593	0.00	0.166	0.0098
52	4-2000	East Branch Black River at Elvira	211	1923-36	4174	533	0.07	0.107	0.0086
55.9	4-2014	West Branch Rocky River at West View	146	1931, 1960 1962-63	4132	576	0.10	0.035	0.0152
56	4-2015	Rocky River near Berea	269	1924-35; 1944-63	4244	586	0.40	0.082	0.0171
67.5	4-2072	Tinkers Creek at Bedford	66.6	1946-63 1962-63	4201	763	0.60	0.0762	0.039
71	4-2090	Chagrin River at Willoughby	231	1928-36; 1940-63	4313	606	0.25	0.104	0.044
72	4-2095	Grand River near North Bristol	89.7	1943-47	435.2	672	0.35	0.043	0.0292
73	4-2100	Phelps Creek near Windsor	26.4	1943-56	432.3	791	0.72	0.0223	0.0142
75	4-2105	Grand River near Rome	276	1943-47	4293	663	0.00	0.0316	0.0216
76	4-2120	Grand River near Madison	387	1923-35; 1939-63	4448	707	0.10	0.166	0.0086
80	4-2130	Connect Creek at Connect	176	1923-35; 1931-63	4239	631	0.70	0.036	0.0170

Low Flow Frequency And Storage
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